The association of natural hydrogen and nitrogen : the ammonium clue ?

Insights from geochemical modeling

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Motivation of the study

Seepage of H2… and N2

- Hydrogen systems of the Oman and New Caledonian ophiolite and of the Brazil's Sao Francisco Basin :
	- Seep of H2 associated with N2 at the central part of the hydrogen systems
	- Seep of (almost) pure N2 at the periphery of the hydrogen systems

Ammonium (NH4+) as the source of N2 ?

- NH4+ (ammonium)-bearing natural minerals are reported in the literature
- NH3 (ammonia; the conjugate base of NH4+) is known to decompose into N2 and H2 with temperature :

 2 NH3 = N2 + 3 H2

- In the lower crust, at temperature around 250 °C, an isotopic study of Bebout et al. (1992) demonstrated that nitrogen generation was linked to ammonium reactivity
- \rightarrow Does NH4+ could be the source/precursor of the emitted N2 in the hydrogen systems ? ⁴

Geochemical models set-up

Initial chemical systems = 'simple' model systems

• **(A) Redox-unbuffered systems**

- Water + FeCl2 $Fe^{2+} + H^+ = Fe^{3+} + 0.5 H_2$
- Water + NH4Cl
- $-$ Water $+$ NH4Cl + FeCl2
- FeCl2 and NH4Cl salts provide ferrous iron & ammonium to the water
- Plus one volume of 1 L possibly hosting a multicomponent gas phase

Initial chemical systems = 'simple' model systems

• **(B) Redox-buffered systems**

- Water + Magnetite-Hematite (*'Control' system)*
- *Water +* Fayalite-Magnetite-Quartz *('Control' system)*
- Water + Magnetite-Hematite + NH4Cl
- Water + Fayalite-Magnetite-Quartz + NH4Cl
- The FMQ & MH mineral assemblages set the oxygen fugacity within the systems ; and provide ferrous & ferric iron sources & sinks
- FMQ: highly reducing conditions; MH: lower ;
- Plus one volume of 1 L possibly hosting a multicomponent gas phase

Initial chemical systems = 'simple' model systems

- **(B) Redox-buffered systems (water-mineral systems)**
	- Water + Fayalite-Magnetite-Quartz + NH4Cl

- Done with the version 3 of the PHREEQC software with the BRGM's Thermoddem database
- Within the 25-300 °C temperature range (Earth surface to ~10 km depth (normal geothermal gradient))
- Considered aqueous, mineral and gaseous species :
	- H2O-Cl-Fe(+II)-Fe(+III)-N(-III)-N(0)-N(+III)-N(+V)-Si
	- Fayalite (Fe2SiO4)
	- Magnetite (Fe3O4)
	- alpha-Quartz (SiO2)
	- Hematite (Fe2O3)
	- H2(g), H2O(g), HCl(g), N2(g), NH3(g), O2(g)

- Based on Law of Mass Action (LMA) resolution
- The K's of formation-dissolution reactions are calculated as function of temperature

$$
\log K_T = A_1 + A_2 T + \frac{A_3}{T} + A_4 \log T + \frac{A_5}{T^2} + A_6 T^2
$$

• The pressure conditions of the calculations are (='convention') :

 $-25-100$ °C : 1 bar

– 100-300 °C : Saturation vapor pressure of water : 1 to 86 bar

- Aqueous solution activity model :
	- B-dot for the charged species (ions)
	- 'Drummond' for dissolved gases (H2(aq), O2(aq), …)
	- Other neutral species : activity of 1
	- H2O(l) (solvent) with a specific equation
- PVTx & gas species activity (fugacity) of the multicomponent gas mixture : Peng-Robinson Equation of State (PR EoS)
	- Inputs : critical temperature & pressure & acentric factors (from phreeqc.dat, NIST and Thermosolver database)
	- 'hard-coded' kij's

$$
\log\gamma_i=-\frac{A\;z_i^2\sqrt{I}}{1+a_i^0B\sqrt{I}}+\dot{BI}
$$

$$
\log\,\gamma_{i(aq)}=-\left(C+FT+\frac{G}{T}\right)\!I+(E+HT)\!\left(\!\frac{I}{I+1}\!\right)
$$

$$
a_{H_2O} = 1 - 0.017 \sum_i^{N_{aq}} \frac{n_i}{W_{aq}}
$$

Results

• Redox-unbuffered systems

0.100 0.008 0.099 0.006 N(-III)_Water+NH4Cl $N = 0.098$
 $N = 0.097$ N(-III)_Water+NH4Cl+FeCl2 H_2 , N_2 (moles) H2_Water+NH4Cl 0.004 H2_Water+NH4Cl+FeCl2 N2_Water+NH4Cl N2 Water+NH4Cl+FeCl2 0.002 0.096 0.095 0.000 R 25 75 125 175 225 275 T (°C)

Water + NH4Cl

 $NH_4^+ = 1.5 H_2 + 0.5 N_2 + H^+$

Redox-buffered systems

 $H2_{TOT}$ & N2_{TOT} (Gas & Water)

- Redox-buffered systems
	- Hydrogen generation
		- Water $+$ Magnetite-Hematite + NH4Cl

 $Fe₃O₄$ (magnetite) + 0.5 $H₂O$ = 1.5 $Fe₂O₃$ (hematite) + 0.5 $H₂$

• Water + Fayalite-Magnetite-Quartz + NH4Cl

1.5 $Fe₂SiO_{4(s)}$ (fayalite) + $H₂O₍₁₎$ = $Fe₃O_{4(s)}$ (magnetite) $+ SiO_{2(s)}$ (quartz) + 0.5 $SiO_{2(aq)}$ (aqueous silica) + $H₂$

$$
Fe^{2+} + H^+ = Fe^{3+} + 0.5\,H_2
$$

- Redox-buffered systems
- Nitrogen generation
	- Water + Magnetite-Hematite + NH4Cl
		- Nitrogen generation occurs within 25-300°C

4 Fe₂O₃ (hematite) + NH_4^+ = 0.5 N₂ + 2.5 Fe₃O₄ (magnetite) $+0.5 \text{ Fe}^{2+} + 2 \text{ H}_2\text{O}$

- Water + Fayalite-Magnetite-Quartz + NH4Cl
	- Nitrogen generation starts at 200 °C

 $2NH_4^+ + 3Fe_3O_{4(s)}$ (magnetite) + 4.5 $SiO_{2(s)}$ (quartz)

 $N_2 + 4 Fe_2SiO_{4(s)}$ (fayalite) + 0.5 $SiO_{2(aq)}$ (aqueous silica) + $Fe^{2+} + 4 H_2O$

$$
3\,\,Fe^{3+} + NH_4{}^+ = 0.5\,\,N_2 + 3\,\,Fe^{2+} + 4\,\,H^+
$$

Typology of the produced gases

• …

17

Implications for hydrogen systems comprehension

$$
(c)
$$

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Outlooks / improvements of the study

To compare with 'hydrothermal' lab experiments

- Ideally on the same 'simple' systems
- Synthetic fluid inclusions in flexible gold cells experiments

To compare with 'hydrothermal' lab experiments

• Coupled Fused Silica Capillary Capsules (FSCCs) and Raman microspectroscopy experiments

• …

• NH4-bearing minerals dissolution

Holloway and Dahlgren (2002)

• K+/NH4+ exchange on clay minerals' surface \rightarrow cation exchange

Appelo and Postma

To investigate T and P conditions beyond 300 °C and 86 bar

Pressure :

PHREEQC is 'designed' to explore 'high' pressures

Figure 5. The solubility of gypsum and anhydrite as a function of temperature at 1, 500, and 1,000 bars. Data points from Blount and Dickson (1973); lines calculated by PHREEQC.

To investigate T and P conditions beyond 300 °C and 86 bar

• Temperature :

- PHREEQC calculations above 300 °C
	- B-dot model : extrapolation of the A(T), B(T) and Bdot(T) parameters to T > 300 °C BUT below the critical T of water (374 $^{\circ}$ C) ?
- Use of other softwares :
	- SupPHREEQC

$$
\log\gamma_i=-\frac{A\;z_i^2\sqrt{I}}{1+a_i^0B\sqrt{I}}+\dot{BI}
$$

• HCh (Shvarov, 2008) (Gibbs Energy Minimization, not Law of Mass Action)

To consider the kinetics of redox and mineral reactions

• Mineral reactions = mineral dissolution-precipitation reactions

Redox reactions kinetics : catalysts, ...?

A page of advertising

- Collaboration with labs (master and PhD students supervision, …)
- Execution of R&D modeling studies (mainly) with PHREEQC
	- 0D geochemical models
	- Reactive Transport (RT) models
- PHREEQC training courses :

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!!! Next PHREEQC course training !!!
             11, 12 and 13 February 2025
     In-person; in (the beautiful town of) Paris
https://geochemical-consulting.com/phreeqc-training-courses
```
Thanks for your attention !

Appendices

• **Redox-unbuffered systems**

Water + FeCl2 Water + NH4Cl +/- FeCl2

$$
Fe^{2+} + H^+ = Fe^{3+} + 0.5 \; H_2
$$

 $NH_4^+ = 1.5 H_2 + 0.5 N_2 + H^+$

• **Redox-buffered systems**

Chemical reactions

+ Aqueous solution

Conclusion

- Gas-water-mineral equilibrium calculations up to 300°C were executed with PHREEQC on initial model systems including ammonium (NH4+)
- \rightarrow To see if this ion could be a 'precursor' in the generation of hydrogen and/or nitrogen seeping out of hydrogen systems
- The below reactions involving NH4+ were evidenced
- Under highly reduced redox conditions (FMQ fO2 buffering) the produced gas is mainly hydrogen ; while under lower reduced redox conditions (MH fO2 buffering), the produced gas is mainly nitrogen
- Implications for hydrogen systems comprehension are addressed

$$
\frac{NH_4^+ = 1.5~H_2 + 0.5~N_2 + H^+}{3~Fe^{3+} + NH_4^+ = 0.5~N_2 + 3~Fe^{2+} + 4~H^+}
$$

```
NH4-feldspar # Buddingtonite
       NH4A]Si3O8 + 4H+ = NH4+ + Al+3 + 3 SiO2 + 2 H2O
       log_k-2.7243-analytic
                      -7.434e1 3.080e-1 0 0 0 -2.270e-4
NH4-muscovite # Tobelite
       NH4A]3Si3O1O(OH)2 + 10 H+ = NH4+ + 3 A1+3 + 3 SiO2 + 6 H2Olog_k6.8109
       -analytic
                      -6.638e1 3.170e-1 0 0 0 -2.386e-4
```